OCS Report MMS 2000-033

EXXON LYDONIA CANYON BLOCK 133 No. 1 WELL

Geological and Operational Summary

Edited by:

Gary M. Edson Donald L. Olson Andrew J. Petty

U. S. Department of the Interior Minerals Management Service Gulf of Mexico OCS Region Office of Resource Evaluation

New Orleans May 2000

CONTENTS

Abbreviations, vi

- Introduction, 1
- Operational Summary, 2
- Well Velocity Profile, 8
- Lithologic Interpretation, 13

Biostratigraphy, 17

- Formation Evaluation, 22
- Geothermal Gradient, 26
- Kerogen Analysis, 28
- Petroleum Geochemistry, 33
- Company-submitted Data, 41
- Selected References, 42

ILLUSTRATIONS

- Figure 1. Map of the North Atlantic offshore area showing well locations, 3
 - 2. Location plat for the Exxon Block 133 No. 1 well on the OCS Lydonia Canyon NK 19-12 protraction diagram, 4
 - 3. Daily drilling progress for the Exxon Lydonia Canyon Block 133 No. 1 well, 5
 - 4. Casing and abandonment diagram for the Exxon Lydonia Canyon Block 133 No. 1 well, 7
 - 5. Well velocity profile for the Exxon Lydonia Canyon Block 133 No. 1 well, plotted against depth, 10
 - 6. Well velocity profile for the Exxon Lydonia Canyon Block 133 No. 1 well, plotted against two-way travel time, 11
 - Columnar chart of the lithology, biostratigraphy, and paleobathymetry of the Exxon Lydonia Canyon Block 133 No. 1 well, 14
 - 8. Well temperatures and geothermal gradient for the Exxon Lydonia Canyon Block 133 No. 1 well, 27
 - 9. Relationships among coal rank, percent R₀, TAI, spore color, and thermal zones of hydrocarbon generation, 29
 - 10. Graph of kerogen types and organic thermal maturity for the Exxon Lydonia Canyon Block 133 No. 1 well, 30
 - Total organic carbon analysis for the Exxon Lydonia Canyon Block 133 No. 1 well, 32
 - 12. Low-molecular-weight hydrocarbon analysis for the Exxon Lydonia Canyon Block 133 No. 1 well, 35
 - Burial diagram for the Exxon Lydonia Canyon Block 133 No. 1 well, 40

TABLES

- Table1. Well statistics, 6
 - 2. Velocity data, 9
 - 3. Well velocity intervals, 12
 - 4. Well logs, 22
 - 5. Well log interpretation summary, 22
 - 6. Sidewall core analysis summary, 23
 - 7. Conventional core analysis summary, 23
 - 8. Drill stem test results, 24
 - 9. Hydrocarbon show summary, 25
 - 10. Total organic carbon, 33
 - 11. Low-molecular-weight and gasoline-range hydrocarbon analyses, 36

API	 American Petroleum Institute
bbl	 barrels
BOP	 Blow out preventer
CNL	 Compensated neutron log
CPI	 Carbon Preference Index
COST	 Continental Offshore Stratigraphic Test
DST	 drill stem test
EQMW	 equivalent mud weight
FDC	 compensated formation density log
FEL	 from east line
FNL	 from north line
FSL	 from south line
FWL	 from west line
k	 permeability
KB	 kelly bushing
LS	 limestone
m	 meter (s)
md	 millidarcy
MYBP	 million years before present
OCS	 Outer Continental Shelf
ppf	 pounds per foot
ppg	 pounds per gallon
ppm	 parts per million
psi	 pounds per square inch
R _O	 vitrinite reflectance
SS	 sandstone
Sw	 water saturation
TAI	 thermal alteration index
TD	 total depth
TIOG	 threshold of intense oil generation
TOC	 total organic carbon
UTM	 Universal Transverse Mercator
φ	 porosity

ABBREVIATIONS

INTRODUCTION

Taken and adapted from MMS published report MMS 89-0007

The Exxon Lydonia Canyon (LC) Block 133 No. 1 well was the first industry wildcat well drilled in Georges Bank Basin. Spudded on July 24, 1981, the well is about halfway between the Continental Offshore Stratigraphic Test (COST) G-1 and G-2 wells. It is about 21 miles eastsoutheast of the COST No. G-1 well. The Exxon LC Block 133 No. 1 well was drilled by a semisubmersible rig in 225 feet of water on the continental shelf about 110 miles east-southeast of Nantucket Island and 35 miles from the shelf edge.

Exxon's primary drilling target was a possible patch reef at a depth of approximately 12,600 feet of expected Middle Jurassic (Callovian) age. Drilling results revealed that the reef was actually a 1,120-foot thick interval of igneous rock interbedded with limestone and minor amounts of sandstone. Hurtubise and others (1987) divided this igneous sequence into an upper, 740-foot thick extrusive unit and a lower, 250-foot thick intrusive unit. The COST No. G-2 well penetrated a similar Middle Jurassic limestone containing up to 5 percent volcaniclastic fragments, mostly green and gray tuff, between 11,720 and 11,900 feet (Amato and Simonis, eds, 1980). There were no significant hydrocarbon shows in the Exxon LC Block 133 No.1 well. Four conventional cores (three recovered) and 158 sidewall cores (72 recovered) were attempted. The single drill stem test recovered only water. Petrophysical and geochemical tests indicate that those depths with sufficient reservoir quality

(above 12,630 feet) do not have sufficient maturity to generate hydrocarbons. Rocks below 12,600 feet show poor reservoir quality, and kerogen types indicate the probability of only gas.

This report relies on geologic and geophysical data provided to the Minerals Management Service (MMS) by Exxon Company, U.S.A. (Exxon) according to Outer Continental Shelf (OCS) regulations and lease stipulations. The data were released to the public after the LC Block 133 lease No. OCS-A-0170 was relinquished on December 30, 1982. Interpretations of the data contained in this report are those of MMS, unless otherwise stated, and may differ from those of Exxon. Depths are measured from kelly bushing unless otherwise indicated.

The information contained herein is taken from MMS OCS Report MMS 89-0007, published in 1989 and now out of print. No attempt has been made to provide more recent geologic, geochemical, or geophysical interpretations or data, published or unpublished, since the initial report was issued.

This report is initially released on the MMS Internet site

http://www.gomr.mms.gov, and, together with the other Georges Bank well reports, on a single compact disk (CD). At a later date, additional technical data, including well "electric" logs will be added to the CD.

OPERATIONAL SUMMARY

Taken and adapted from K. U. Siddiqui, MMS published report MMS 89-0007

The Exxon Lydonia Canyon (LC) Block 133 No. 1 well (figures 1 and 2) was drilled by the North Star Drilling Company's *Alaskan Star* semisubmersible drilling rig for Exxon Company, U.S.A. The well was spudded on July 24, 1981 in 225 feet of water. Daily drilling progress is shown for the well in figure 3, and well statistics are presented in table 1. The primary geologic objective was a possible Middle Jurassic (Callovian) patch reef at approximately 12,600 feet.

Five strings of casing were set in the well (figure 4). The 30-inch casing was set to 550 feet with 700 sacks of cement; the 16-inch casing was set to 993 feet with 1,625 sacks of cement; the 13 3/8-inch casing was set to 4,100 feet with 1,200 sacks of cement; the 9 5/8-inch casing was set to 11,754 feet with 1,150 sacks of cement; the 7-inch casing was set to 14,050 feet with 615 sacks of cement. Class H cement was used for all casings.

Lignosulfonate, freshwater, and seawater muds with an average weight of 9.3 pounds per gallon (ppg) were used as drilling fluids to a depth of 1,050 feet. The mud weight was raised to 9.4 ppg at a depth of 3,833 feet and to 9.8 ppg at 10,878 feet. Mud weight reached 11.3 ppg at 13,525 feet and remained at that weight to the total depth of 14,118 feet. Mud viscosity averaged 52 seconds, fluctuating between 41 and 68 seconds, in the first 1,050 feet and averaged about 48 seconds for the remainder of the well. Mud pH averaged 11.1 with minor fluctuation for the entire well. Chloride concentrations began at 450 parts per million (ppm), increased to 5,800 ppm at 12,578 feet, and dropped to 2,600 ppm at TD.

Abandonment procedures are also shown in figure 4. The 7-inch casing was perforated at 12,392 to 12,408 feet and 12,418 to 12,430 feet and a retainer was set at 12,275 feet. A plug was set from 12,225 to 12,586 feet with 50 sacks of cement and was tested at a pressure of 2,500 pounds psi for 15 minutes. The 7inch casing was cut at 4,249 feet, and the second plug was set between 4,125 and 4,400 feet with 75 sacks of cement and tested at a pressure of 1,000 psi for 15 minutes. The 9 5/8-inch casing was perforated from 3,898 to 3,900 feet, a retainer set at 3,840 feet, and a 150-sack plug was set at 3,780 to 4,334 (annular) feet and tested at a pressure of 1,000 psi for 15 minutes. The 9 5/8-inch casing was cut at 1,150 feet. A 150-sack plug was set from 1,025 to 1,300 feet. The 13 3/8-inch casing was perforated at 796 to 798 feet. a retainer was set at 716 feet, and a 200-sack plug was set from 716 to 1,197 (annular) feet and tested at a pressure of 1,000 psi for 15 minutes. The 13 3/8-inch casing was cut at 558 feet. The 16-inch casing was cut at 335 feet, and the 30-inch casing was cut at 327 feet. The surface plug was set at 420 to 710 feet with 275 sacks of cement and pressure tested.



Figure 1. Map of the North Atlantic offshore area showing well locations. The Exxon Lydonia Canyon Block 133 No. 1 well is highlighted in red.. Bathymetry is in meters.

N X X	128 Protraction Diagram NK 19 - 11	89 Protraction Diagram NK 19 - 12	90
	172	8,852.61' FNL 133 4,343.82' FEL 40° 49' 05.230" N Lat. 67° 56' 03.484" W Long	- 134
	216	177	178
	0	8,000' 16,	000'

Figure 2. Location plat of the Exxon Block 133 No. 1 well on the OCS Lydonia Canyon NK 19-12 protraction diagram..



Figure 3. Daily drilling progress for the Exxon Lydonia Canyon Block 133 No. 1 well.

Well Identification:	API #61-040-00002 Lease No. OCS-A-0170
Surface location:	Lydonia Canyon NK 19-12 Block 133 8,852.61' FNL 4,343.82' FEL
	Latitude: 40° 49' 05.230" N Longitude: 67° 56' 03.484" W
	UTM coordinates: X = 589,876.00m Y = 4,518,901.72m
Bottomhole location:	167.8 feet N and 129.1 feet E of surface location
Proposed total depth:	15,500 feet
True vertical depth:	14,110 feet
Measured depth:	14,118 feet
Kelly bushing elevation:	85 feet
Water depth:	225 feet
Spud date:	July 24, 1981
Reached TD:	October 20,1981
Off location:	November 24, 1981
Final well status:	Plugged and abandoned

Table 1. Well statistics

Note: All depths indicated in this report are measured from the kelly bushing, unless otherwise indicated. Mean sea level is the datum for the water depth.

The blowout preventer and guide base were pulled, the *Alaskan Star* moved off location on November 24, 1981, and a post-abandonment, seafloor site survey was performed by John Chance and Associates.

	Depth Below KB	(Feet)
	Sea Level	85
Se	eafloor	310
30) Inch Casing Cut	327
	δ Inch Casing Cut	335
Τα	op of Cement	420
30 13 3/8	Inch Casing Set Inch Casing Cut	550 558
Bottor Retair	n of Cement her, Top of Cement	710 716
13 3/8	Inch Casing Perforated	796 - 798
16 Inc. Top of Cen 9 5/8 Inch Bottom of Top of Cen Retainer 9 5/8 Casin 13 3/8 Inch Top of Cement 7 Inch Casing 0 Bottom of Cem Bottom of Cem Bottom of Cem Bottom of Cem	h Casing Set nent Casing Cut Cement Cement nent ng Perforated n Casing Set Cut lent ent ent mg Set	993 1,025 1,150 (annular) 1,300 3,780 3,840 3,898- 3,900 4,100 4,125 4,249 4,334 (annular) 4,400 11,754 12,225 12,275 12,275
Bottom of Cement		12,586
		13,000
Top of Cement 7 Inch Casing Set		13,621 14,050
Total Depth		14,118

Figure 4. Casing and abandonment diagram for the Exxon Lydonia Canyon Block 133 No. 1 well.

WELL VELOCITY PROFILE

Taken and adapted from A. O. Tanner, MMS published report MMS 89-0007

Velocity surveys of the Exxon LC Block 133 No. 1 well were conducted by Seismograph Service Corporation (Birdwell Division) and Schlumberger, Ltd. In addition to a sonic log, Schlumberger ran a velocity checkshot survey between 8,917 and 13,551 feet, whereas Birdwell conducted a velocity checkshot between 4,817 and 11,657 feet. The checkshot data, together with that for the other nine wells drilled on Georges Bank, were given by MMS to Velocity Databank, Inc. at their request after all leases had been relinquished or had expired. Velocity Databank calculated interval, average, and RMS velocities, plotted time-depth curves, and tabulated the data. The Birdwell data were used for the Exxon LC 133 No. 1 well. Table 2 presents well depth, two-way travel time, and the calculated velocities. Figures 5 and 6 show interval velocity, average velocity, and RMS velocity plotted against depth and against two-way travel time. These depths are relative to sea level.

A lithology column is also shown in figure 5, and velocity intervals are given in table 3. The first checkshot at 227 feet is near the seafloor and yields water-column velocities. From 4,817 feet (the first checkshot below mudline) to 11,657 feet (the deepest data), six intervals are identified based on changes in relative interval velocities.

Interval I This interval contains the first two data points, including the water column and sediment to 5,000 feet. The low to moderate velocities increase with depth, reflecting the progression from water to unconsolidated sediment to

lithified siliciclastic rock. The velocity increase is also a consequence of greater rock densities with increasing depth, an effect that continues through subsequent intervals. This interval is Lower Cretaceous to Tertiary.

Interval II This interval is identified on the basis of intermediate interval velocities reflecting limestone, sandstone, shale, and siltstone. The interval is Lower Cretaceous.

Interval III This interval is identified on the basis of intermediate interval velocities that reflect increasing finer grained siliciclastic sediments and less limestone with depth. The interval is Middle and Upper Jurassic.

Interval IV This interval is identified on the basis of an increase in interval velocities. The lithologic description (and mud log show interbedded siliciclastics and calcareous siliciclastics throughout the interval. The interval is Middle Jurassic.

Interval V This interval is identified on the basis of high velocities, averaging 17,280 feet per second, which correlate with limestone and siliciclastic interbeds. This interval is Middle Jurassic.

Interval VI The final interval includes a single reading, the deepest in the well at 11,657 feet, with an extremely high interval velocity of 40,002 feet per second. This velocity is probably in error, being in excess of that for any crustal rock type.

Depth	Two-way Travel	Interval	Average	RMS Velocity
(Feet)	Time	Velocity	Velocity	(Feet/Sec.)
	(Seconds)	(Feet/Sec.)	(Feet/Sec.)	
227	0.090	5,044	5,044	5,044
4,817	1.340	7,343	7,189	7,211
5,217	1.426	9,302	7,316	7,354
5,417	1.466	10,000	7,390	7,439
5,617	1.498	12,499	7,499	7,582
5,817	1.534	11,111	7,584	7,684
6,017	1.572	10,526	7,655	7,765
6,217	1.612	10,000	7,713	7,828
6,417	1.646	11,764	7,797	7,929
6,617	1.674	14,285	7,905	8,076
7,017	1.744	11,428	8,047	8,237
7,217	1.774	13,333	8,136	8,349
7,417	1.804	13,333	8,222	8,456
7,817	1.874	11,428	8,342	8,586
8,017	1.902	14,285	8,430	8,697
8,217	1.934	12,499	8,497	8,773
8,417	1.966	12,500	8,562	8,846
8,617	2.002	11,111	8,608	8,892
8,817	2.030	14,285	8,686	8,988
9,017	2.054	16,666	8,779	9,116
9,217	2.090	11,111	8,820	9,154
9,417	2.114	16,666	8,909	9,273
9,617	2.148	11,764	8,954	9,318
9,817	2.174	15,384	9,031	9,413
10,017	2.208	11,764	9,073	9,454
10,217	2.230	18,181	9,163	9,579
10,417	2.256	15,384	9,234	9,666
10,617	2.288	12,499	9,280	9,711
10,817	2.318	13,333	9,333	9,766
11,017	2.344	15,384	9,400	9,846
11,217	2.362	22,222	9,497	9,999
11,417	2.392	13,333	9,545	10,047
11,617	2.414	18,181	9,624	10,151
11,657	2.416	40,002	9,649	10,212

 Table 2. Velocity data



Figure 5. Well velocity profile for the Exxon Lydonia Canyon Block 133 No. 1 well, plotted against depth, with biostratigraphic ages and generalized lithologies. Intervals are explained in text.



Figure 6. Well velocity profile for the Exxon Lydonia Canyon Block 133 No. 1 well, plotted against two-way travel time.

Interval	Depth Range	Interval Velocity	Average Interval Velocity
	(feet)	Range (feet/second)	(feet/second)
Ι	0- 5,000	5,044- 7,343	6,194
II	5,000- 6,500	9,302-12,499	10,743
III	6,500- 8,700	11,111-14,285	12,689
IV	8,700-11,000	11,111-18,181	14,276
V	11,000-11,630	13,333-22,222	17,280
VI	11,630-11,657	40,002	40,002

 Table 3. Well velocity intervals

Well cuttings indicate limestone at 11,657 feet; igneous rock fragments were not recognized until about 12,500 feet. At about 65 microseconds per foot, the sonic log does not indicate an anomalously high velocity that corresponds to the deepest checkshot velocity. Altogether, well logs are consistent with interbedded carbonates and siliciclastics below 11,000 feet. Interval VI is Middle Jurassic.

LITHOLOGIC INTERPRETATION

Taken and adapted from A. C. Giordano, MMS published report MMS 89-0007

These descriptions are based on well cutting samples, collected at 10-foot intervals from 630 to 14,118 feet in the Exxon LC Block 133 No. 1 well. Sample quality ranged from fair to good, depending on the amount of cavings and the degree of washing. Additional lithologic control was provided by the physical formation "mud" log, 72 sidewall core analyses, descriptions of three conventional cores, and examination of chips from the conventional cores. Depths of lithologic boundaries are adjusted with reference to "electric" and "mud" logs. All depths are from kelly bushing. Rocks penetrated are divided into gross lithologic-stratigraphic units, and a lithologic column appears in figure 7.

LITHOLOGIC DESCRIPTIONS

From 630 to 1,180 feet, the section consists of gray to gray-green, soft, slightly silty claystone. A glauconitic clay, between 1,180 and 1,280 feet, marks an unconformity between the Paleocene and oxidized Upper Cretaceous (Maestrichtian) rocks.

Between 1,280 and 1,950 feet, the section consists of light-gray to gray-green, slightly calcareous claystone and argillaceous siltstone. Glauconite, fragments of megafossils, foraminifera, carbonaceous plant fragments, fine mica, and pyrite also occur.

Between 1,950 and 2,720 feet, the section consists of unconsolidated coarse quartz grains and granules, interbedded with light-gray, slightly calcareous silty claystone and clayey siltstone. Fossil fragments and inertinite occur within the clays. The dominant lithology between 2,720 and 2,960 feet is light-gray to white, microcrystalline, fossiliferous limestone that is mottled in places.

Between 2,960 and 3,800 feet the section consists of sand, shell fragments, glauconite pellets, mica flakes, thin streaks of lignite, and calcite cement. The sand is quartz, fine to granular, subangular to subrounded, and unconsolidated. Carbonaceous, dark-gray shales also occur. From 3,800 to 4,880 feet, a red-to-gray mudstone is interbedded with the sandstone. The sandstone in this mudstone sequence is similar to that of the overlying sandy section, but also includes gray, fine-to-medium-grained, muddy, friable wackestone.

The section from 4,880 to 6,610 feet consists of limestone, sandstone, siltstone, mudstone, and shale. The upper portion of this section, from 4,880 to 5,600 feet, is cryptocrystalline, sandy limestone and glauconitic, pyritic shale and siltstone interbedded with glauconitic and fossiliferous, calcareous sandstone. From 5,600 to 6,610 feet, thick-bedded sandstone with shale and coal interbeds forms the lower portion of this section. Much of the lower sandstone is poorly consolidated and very coarse to medium in texture. Present also is very fine-tomedium grained, subangular, moderately well-sorted sandstone tightly cemented by calcite and silica. Between 5,520 and

	DEPTH IN FEET		AGE	ГІТНОГОСУ	PALYNOLOGY	NANNOFOSSIL	FORAMINIFERA	NON MARINE	P. 0 5	ALEOE 50' 30 SHEL MIDDLE	NVIRC 00' 60 F OUTER	ONMEN 0' 15 SLOF UPPER	IT 00' 2E LOWER
16"	-	ARY	Miocene		(730)								-
Casing at 993'	- — 1000 -	TERTI	Indeterminate Age		(1,100)			ſ					-
	-		Early Eocene-Paleocene Maestrichtian		(1 280)		(1,310)	~~~~	\vdash			~~~~	~~~~
			Campanian		(1,200)	(1,340)							-
	-	SUC	Santonian			(1,460)							-
	- - 2000	ETACEC	Coniacian				(1,820)						-
	- - - - 2500 -	LATE CRE	Turonian			(2,090)							- - - -
	- -		Cenomanian		(2,720)	(2,720)	(2,660)						-
	- 3000 - -				(2,900)								-



Figure 7. Columnar chart of the lithology, biostratigraphy, and paleobathymetry of the Exxon Lydonia Canyon Block 133 No. 1 well. Lithologic interpretations from examination of cuttings; lithologic breaks picked from well logs. Within columns, depths refer to uppermost occurrence of index fossils listed in Biostratigraphy chapter. Stage tops based on paleontology. Biostratigraphy and bathymetric interpretations become less reliable with increasing depth.

6,120 feet, numerous thin beds of coal occur in dark, carbonaceous shale, browngray shale, and light-gray mudstone. From 6,350 to 6,610 feet, the sandstone is interbedded with limestones and shale.

The section from 6,610 to 7,250 feet consists of sandstone, shale, and limestone. The sandstone beds are sparse and thin. The shales are gray to dark gray, silty, and calcareous. The limestone is light gray to tan, cryptocrystalline, oolitic, fossiliferous, and argillaceous. Contact with the underlying sandy unit is gradational.

The section between 7,250 and 10,850 feet contains sandstone, red and gray shale, and streaks of coal. Most of the sandstone is white, fine to medium grained, subangular, moderately to well sorted, and friable to tightly cemented by calcite. Muscovite, biotite, and chlorite are common. Some sandstone beds in the lower part of the sequence are iron stained. The shales from 7,250 to 8,960 feet are gray to dark gray, silty, and variably calcareous. Red shales appear below 8,960 feet.

The section from 10,850 to 12,390 feet consists mainly of oolitic packstone. Present also are fine-grained, well-sorted sandstones, dark-gray, argillaceous shales, and gray mudstones.

The interval from 12,390 to 13,130 feet is primarily an extrusive unit consisting of volcaniclastic rock and basalt flows with minor limestone interbeds (Hurtubise and others, 1987). Sandstone interbeds also occur in the uppermost part of this unit. Below 12,580 feet, tuff comprises as much as 5 percent of each sample and persists in trace amounts to 13,260 feet. The tuff is green and gray, vesicular, devitrified glass and occurs as inclusions.

From 13,130 to 14,118 feet (TD) the section consists of tight, cryptocrystalline limestones, within which an intrusive basalt occurs from 13,260 to 13,510 feet. The basalt is very fine grained with a dark gray matrix containing scattered crystals of quartz, olivine, calcite, and chlorite. Fractures are filled with veins of soft serpentine. A radiometric date of 136 million years indicates an Early Cretaceous age (Hurtubise and others, 1987).

POTENTIAL RESERVOIR ROCKS

Sidewall core analysis for the sandstones above 6,610 feet shows a porosity range from 17 to 25 percent and permeability as high as 68 millidarcies (md). Most samples are calcite cemented. Cementation appears to have occurred relatively early, before significant compaction. Primary porosity is preserved in many of the better sorted sands.

Reservoir quality varies from 6,610 to 10,850 feet. No porosity is evident from 6,650 to 8,080 feet. From 8,080 to 9,830 feet, visually estimated porosity is from 15 to 23 percent. From 9,900 to 10,850 feet, it is 8 to 14 percent. Silty sandstones and shales are either compacted or tightly cemented with dolomite or calcite.

From 10,850 to 14,118 feet, visual porosity ranges from negligible to 15 percent. Sandstone in sidewall cores from 12,000 to 14,007 feet has indicated porosities of 4.5 to 17 percent but very low permeabilities, ranging from <0.1 to 3.4 md. The limestone beds are very tight, and no significant secondary porosity is evident. Siltstones are tightly cemented with silica and calcite. Three conventional cores recovered from intervals between

12,250 and 13,491 feet contain cryptocrystalline limestone, shales, and the intrusive basalt, none with significant porosity.

BIOSTRATIGRAPHY

Taken and adapted from H. L. Cousminer, W. E. Steinkraus, and L. E. Bielak, MMS published report MMS 89-0007

Two factors limit the reliability of paleontologic data from exploration wells. (1) Analyses are made from drill cuttings that are often heavily contaminated by cavings from higher in the drill hole. For this reason, only "tops," or the uppermost (last) appearances of species, are recorded. (2) Reworked, older fossil assemblages and individual specimens are commonly reincorporated in detrital sedimentary rocks. These fossils must be recognized so those intervals are not dated older than they really are. In addition, in U.S. offshore Atlantic wells, biostratigraphic control is poor in pre-Late Jurassic strata. Calcareous nannofossils and foraminifera are sparse. Palynomorphs are more common, but their biostratigraphic distribution is not fully documented with reference to the European type-stage localities.

Microfossils were examined from 630 to 14,118 feet (TD) (figure 7) in the Exxon LC Block 133 No. 1 well. The analysis included 128 palynomorph slides prepared from composite 90-foot intervals from the entire range of well sample depths; 266 nannofossil slides representing 30-foot intervals from 730 to 8,620 feet; and 74 foraminiferal slides from 30-foot intervals from 730 to 3,000 feet. The interpreted ages range from Miocene to Bajocian. An unconformity was observed at 1,180 feet.

Planktonic foraminiferal age markers are identified in only three Late Cretaceous samples. Nannofossil markers range in age from Middle Jurassic (Callovian) to Late Cretaceous (Campanian) age. Palynomorphs (dinocysts, spores, and pollen) range in age from Early Jurassic (Sinemurian-Hettangian) to Miocene. However, all palynomorphs older than Middle Jurassic in age are considered to have been reworked in the section. After completion of the initial study, which was based solely on cuttings samples, core samples from 12,258 to 12,326 feet were received from the operator. Palynologic analysis of an additional 20 slides representing nine sample splits within this cored interval indicates it to be of Middle Jurassic age (Bajocian). Environments of deposition ranged from nonmarine to inner shelf.

CENOZOIC

TERTIARY

Miocene (730-820 feet)

This interval contains palynomorph species of <u>Carya</u>, <u>Lycopodium</u>, <u>Osmunda</u>, <u>Quercus</u>, <u>Piceapollenites</u>, <u>Pinuspollenites</u>, <u>Polypodium</u>, and <u>Sphagnum</u>. The absence of <u>Compositae</u>, <u>Chenopodiaceae</u>, and <u>Gramineae</u> indicates a probable midearly Miocene age. The environment of deposition is inner shelf (0-50 feet).

Indeterminate (820-1,100 feet)

Samples from 820 to 1,100 feet are barren. The environment of deposition is probably nonmarine.

Early Eocene-Paleocene (1,100-1,180 feet)

The highest occurrence of Alnipollenites sp. and Laevigatosporites haardtii at 1,100 feet indicates an early Eocene age (Bebout, 1980). The highest occurrence of Stereisporites and Sapotaceoidaepollenites at 1,100 feet indicates Paleocene. Therefore, this interval is considered to be Paleocene to early Eocene in age. Other palynomorphs present include Alnipollenites, Laevigatosporites haardtii, Momipites, Nymphaceae, Sapotaceoidaepollenites, Stereisporites, and Tetrasporites. No nannofossils or foraminifera are present. The environment of deposition is nonmarine.

MESOZOIC

CRETACEOUS

Late Cretaceous

Maestrichtian (1,180-1,340 feet)

The interval from 1,180 to 1,280 feet is barren. However, organic residues match those from below 1,280 feet; therefore, the interval is probably of Late Cretaceous age. The highest occurrence of <u>Lejeunia koslowskii</u> at 1,280 feet indicates a Maestrichtian age (Gorka, 1963; Kjellstrom, 1973). The planktonic foraminifera <u>Globotruncana arca</u> and <u>G</u>. <u>marginata</u> found in this interval support a Maestrictian age. The environment of deposition is inner shelf. The sudden change from nonmarine to inner shelf at 1,180 feet indicates an unconformity.

Campanian (1,340-1,460 feet)

The nannofossil <u>Gartnerago obliquum</u> and the pollen <u>Retitricolpites</u> sp. "M" Wolfe indicate a Campanian age. According to Bebout (1981), the latter does not range above Campanian in the Mid-Atlantic OCS. The environment of deposition is inner shelf.

Santonian (1,460-1,820 feet)

The highest occurrence of the nannofossil <u>Marthasterites furcatus</u> at 1,460 feet marks the top of the Santonian. The environment of deposition is inner shelf but the presence of <u>Pediastrum</u> sp. may indicate nearshore brackish conditions.

Coniacian (1,820-2,090 feet)

The highest occurrence of <u>Globotruncana</u> renzi at 1,820 feet indicates a Coniacian age. The environment of deposition is inner shelf.

Turonian (2,090-2,660 feet)

The Turonian nannofossil <u>Corollithion</u> <u>achylosum</u> was identified at 2,090 feet. The environment of deposition is inner shelf.

Cenomanian (2,660-2,900 feet)

The highest occurrence of the planktonic foraminifera <u>Rotalipora appenninica</u> and <u>R. cushmani</u> at 2,660 feet and the nannofossil C<u>orollithion signum</u> "A" at 2,720 feet indicate a Cenomanian age for this interval. Abundant dinocysts from 2,720 feet to 2,810 feet indicate the upper Cenomanian Cleistrosphaeridium <u>huguonioti</u> subzone of Clarke and Verdier (1967). The environment of deposition is inner shelf.

Early Cretaceous

Albian (2,900 to 3,800 feet)

The highest occurrence of the dinocysts <u>Apteodinium maculatum</u> and <u>Astrocysta</u> <u>cretacea</u> at 2,900 feet mark the top of the Albian. <u>Pediastrum</u> sp., possibly indicating brackish nearshore conditions, is also present at 3,260 feet. <u>Trilobosporites apiverrucatus</u>, from 3,270 feet, does not range above the Albian (Bebout, 1980). The environment is nonmarine to inner shelf for the lower part of the interval.

Aptian (3,800 to 4,880 feet)

Abundant dinocysts at 3,800 feet indicate the <u>Cyclonephelium attadalicum</u> zone, which is confined to the early Aptian on the Scotian Shelf (Williams, 1977). The nannofossils <u>Nannoconus globulus</u> (4,580 feet) and <u>N. wassalli</u> (4,820 feet), as well as Bebout's (1981) spore species 110 (4,700 feet), support an Aptian age for this interval. The spore does not range above the Aptian in the Mid-Atlantic OCS. The environment of deposition is inner to middle shelf.

Barremian (4,880-5,450 feet)

The <u>Tenua anaphrissa</u> subzone at 4,880 feet (Williams, 1977) indicates a Barremian age in the Scotian Shelf-Grand Banks region. Other dinocyst species present that do not range into the Aptian are <u>Muderongia simplex</u>, Polystephanephorus sarjeantii, and <u>Pseudoceratium pelliferum</u> (Williams, 1977). The environment of deposition is inner to middle shelf, shallowing to inner shelf at approximately 5,000 feet.

Hauterivian-Vananginian (5,450-6,340 feet)

The highest occurrence of Bebout's <u>Trilobosporites</u> 132 and 133 indicate a Hauterivian or Vananginian age for the interval from 5,450 to 5,980 feet. These forms do not range above the Hauterivian in the Mid-Atlantic Outer Continental Shelf. The environment of deposition is inner shelf for the upper part of the interval, shallowing progressively to nonmarine in the lower part of the section.

Berriasian (6,340-6,520? feet)

The dinocysts <u>Chlamydophorella</u> sp. "A" Davey, <u>Phoberocysta neocomica</u>, <u>Kleithriasphaeridium</u> sp. "A" Davey, and Bebout's spores 139 and 140 are present in this interval. In England and northwest Europe, <u>Kleithriasphaeridium</u> sp. "A" Davey does not range above the Ryazanian (Berriasian) Speeton Clay (Davey, 1979). Bebout (1981) reports that the spore types 139 and 140 do not range above the Berriasian in the Mid-Atlantic OCS. The environment of deposition is nonmarine.

JURASSIC

Late Jurassic

Tithonian (6,520?-6,970 feet)

A barren zone, 6,520 to 6,610 feet, is oxidized and appears to match the

immediately underlying interval. <u>Ctenidodinium panneum</u> present below 6,610 feet indicates a Tithonian age. The Scotian Shelf <u>Ctenidodinium panneum</u> Zone is provisionally dated as Tithonian (Williams, 1977). Davey (1979) restricted the upper range of this species to lower Tithonian in northwest Europe. The environment of deposition is inner shelf.

Kimmeridgian (6,970-7,330 feet)

The Kimmeridgian top is marked by an abundant dinocyst assemblage including <u>Gonyaulacysta longicornis</u>, <u>G</u>. <u>cladophora</u>, <u>G</u>. <u>mamillifera</u>, <u>G</u>. sp. "H" (Gitmez and Sarjeant, 1972), and <u>Hexagonifera</u> (<u>Senoniasphaera</u>) <u>jurassica</u>. The environment of deposition is inner shelf, shallowing to marginal marine deeper in the section.

Oxfordian (7,330-7,870 feet)

The highest occurrence of <u>Adnatosphaeridium aemulum</u> and <u>Surculosphaeridium vestitum</u> at 7,330 feet indicates the Oxfordian age. The nannofossil species <u>Watznaueria</u> <u>reinhardti</u> and <u>Stephanolithion bigoti</u> were recovered at 7,600 feet. <u>S. vestitum</u> does not range above Oxfordian (Sarjeant, 1979). <u>A. aemulum</u> and <u>S. vestitum</u> do not range above the Oxfordian (Bujak and Williams, 1977). In the Scotian Shelf-Grand Banks region <u>S. bigoti</u> ranges from Oxfordian to lower Kimmeridgian. The environment of deposition is inner shelf.

Middle Jurassic

<u>Callovian (7,870-8,590 feet)</u> The highest occurrence of <u>Stephanolithion spectiosumoctum</u> at 7,870 feet marks the Callovian. The dinocysts <u>Valensiella ovulum</u> and <u>Gonyaulacysta aldorfensis</u> both have their highest occurrences in this interval and do not range above the Callovian in the Scotian Shelf-Grand Banks region (Bujak and Williams, 1977). Abundant also in this interval are the longer-ranging <u>Adnatosphaeridium</u> sp., <u>A. aemulum</u>, <u>Ctenidodinium ornatum</u>, and <u>Sentusidinium</u> spp. The environment of deposition is inner shelf, shallowing to marginal marine in the lower 200 feet.

Bathonian (8,590-9,760 feet)

The highest occurrence of the dinocyst species Gonyaulacysta filapicata at 8,590 feet marks the top of the Bathonian (Bujak and Williams, 1977). Other dinocyst species present include Gonyaulacysta aldorfensis, Sentusidinium spp., Lithodinia jurassica, and Adnatosphaerieium callveri, all of which range above the Bathonian. Present also are Ctenidodinium continuum, C. ornatum, C. pachydermum, Leptodinium spp., and Vanensiella spp. Gocht (1970) described a similar assemblage from the type Bathonian of northwest Germany. A barren interval from 9,480 to 9,760 feet contains detrital minerals and tracheal detritus, probably representing a regressive facies. The environment of deposition is marginal marine to nonmarine.

Bajocian (9,760-13,130 feet)

The highest occurrence of sparse <u>Mancodinium semitabulatum</u> and <u>Mendicodinium reticulatum</u> mark the top of the Bajocian. Present also is abundant Gonyaulacysta filapicata from 9,760 to 10,570 feet. Although numerous palynomorphs that are restricted to Liassic age in several Scotian Shelf-Grand Banks wells (Bujak and Williams, 1977) are present in cuttings from 10,570 to 12,326 feet, these are all considered to be reworked. Core samples from 12.258 to 12,326 feet contain abundant Middle Jurassic (Bathonian-Bajocian) dinoflagellate species, including Adnatosphaeridium callyeri, Ctenidodinium spp., Dichatogonyaulax aff. staruomatos, Gonyaulacysta pectinigera, and Mancodinium semitabulatum. These indicate that the older (Hettangian-Pleinsbachian) species present in both the cuttings and core material are reworked elements. Such species include Corollina meyeriana, Convolutispora klukiforma, Cycadopites subgranulosus, Kraeuselisporites

reissingeri, and Verrucosisporites cheneyi (Bujak and Williams, 1977). According to Hurtubise and others (1987), these palynomorphs represent erosional detritus accumulated during Liassic subaerial exposure and recycled during the subsequent Middle Jurassic marine transgression. The environment of deposition is nonmarine in the upper 500 feet, deepening to inner shelf from approximately 10,200 to 11,500 feet. From 11,500 to approximately 12,600 feet the environment of deposition is marginal marine to inner shelf, becoming nonmarine to 13,130 feet.

Barren (13,130-14,118 feet)

Samples examined from 13,130 to 14,118 feet were barren of fossil material. The environment of deposition was probably nonmarine.

FORMATION EVALUATION

Taken from R. R. Nichols, MMS published report MMS 89-0007

Schlumberger Ltd. ran the following geophysical "electric" logs in the Exxon LC Block 133 No. 1 well to provide

information for stratigraphic correlation and for evaluation of formation fluids, porosity, and lithology:

Table 4. Well logs

Log Type	Depth Interval
	(feet) below KB
DISFL (dual induction-spherically focused log)	997-14,118
GRS (gamma ray-sonic log)	997-14,118
GRN (gamma ray-neutron log)	367-13,659
CNL/FDC (compensated neutron log/compensated formation density log)	4,101-14,118
FDC (compensated formation density log)	4,101-14,118
FIL (fracture identification log)	11,739-14,114

Exploration Logging, Inc. provided a formation evaluation log "mud log" that included a rate of penetration curve, sample descriptions, and a graphic presentation of hydrocarbon shows encountered (550 to 14,115 feet).

The electric logs, together with the mud log and other available data, were analyzed in detail to determine the thickness of potential reservoirs, average porosities,

and feet of hydrocarbons present. Reservoir rocks with porosities less than 5 percent were disregarded. A combination of logs was used in the analysis, but a detailed lithologic and reservoir property determination from samples, conventional cores, and sidewall cores, in addition to full consideration of any test results, is necessary to confirm the estimates shown in table 5.

Depth Interval	Potential Reservoir ¹	Average	SW	Feet of
(feet)	(feet)	Porosity (%)	(%)	Hydrocarbon
1,953-2,038	84	35	NC^2	
2,744-4,260	797	35	NC	
4,782-6,668	1,058	23-33	NC	
6,668-8,030			NC	
8,030-8,974	340	15-24	NC	
9,082-9,867	205	12-18	NC	
9,992-10,130	69	15	NC	
10,220-10,256	34	14	NC	
10,550-10,574	22	9	NC	

Table 5. Well log interpretation summary

Continued

Depth Interval	Potential Reservoir ¹	Average	SW	Feet of
(feet)	(feet)	Porosity (%)	(%)	Hydrocarbon
10,718-10,732	14	13	NC	
10,778-10,788	10	10	NC	
10,834-10,850	16	9	NC	
10,962-11,148	80	8-15	NC	
12,398-12,411	8	23	25	3
12,424-12,434	8	24	26	3
13,464-13,504	37	12 ⁴	10^{4}	4
13,504-14,118			NC	

 Table 5. Well log interpretation summary--continued

¹Generally in beds > 10 feet thick and ϕ > 5%

²NC--not calculated

³Drill stem test of the 12,392-12,408 and 12,418-12,430 foot intervals recovered 564 bbl of water.

⁴From sidewall core analysis at 13,492 and 13,498 feet, $\phi = 12\%$, k = 1 md. Caliper shows severe washout in this zone with corresponding decrease in reliability of electric log responses. No mud log shows were present in this zone. Zone was conventionally cored.

A summary of the sidewall core analysis is shown in table 6.

Table 6. Sidewall core analysis summary

Depth Interval (feet)	Lithology	Porosity Range (%)	Permeability Range (md)
4,785-6,600	Sandstone	17-25	0.2-68
12,000-12,300	Sandstone	4.5-16	<0.1-3.4
12,300-12,800	Sandstone	4.8-17	<0.1-3.0
13,492-14,007	Sandstone	10-13	0.5-1.0

Four conventional cores were taken. Core 3 was not recovered. Lithologic descriptions for cores 1, 2, and 4 were provided by Exxon, but petrophysical core analysis was not available for comparison to specific electric log responses. The conventional core properties are described as follows in table 7.

Table 7.	Conventional	core analys	is summarv

Core No.	Depth Interval (feet)	Lithology	Porosity	Permeability
1	12,250-12,270	Limestone	Tight	Negligible
2	12,272-12,326	Limestone	Tight	Negligible
3 ^a	13,481-13,486			
4	13,488-13,491	Diabase		

^aNot recovered

Table 8 summarizes shows of hydrocarbons encountered in this well, and table 9 lists all hydrocarbon shows. Two drilling breaks were encountered in the interval from 12,400 to 12,435 feet and the mud-logging unit detected significant

 Table 8. Drillstem test results

Test No.	Interval Tested (feet) ¹	Length of Test (hours)	Final Flow Pressure (psi)	Final Shut-in Pressure (psi)	Results/ Recovery
DST-1	12,392-12,430	12.5	8,744	6,770	564 bbl of water, no gas or liquid hydrocarbon, no hydrogen sulfide

¹Drilling breaks occurred in the interval 12,400-12,435 feet with mud log total gas readings increasing from zero to 95 units.

amounts of gas. The mud log total gas readings increased from a background of zero to a high of 95 units (methane only). Well log interpretation in the zone of interest indicates eight feet of potential reservoir between 12,398 and 12,411 feet with an average porosity of 23 percent and a water saturation of 25 percent. An additional eight feet of potential reservoir exists between 12,424 and 12,434 feet with an average porosity of 24 percent and a water saturation of 26 percent. Density log porosity exceeding the sonic log porosity indicates the probability of secondary porosity in the zone of interest. Characteristic crossover responses of the CNL and FDC curves suggest the presence of gas in this interval, supporting the mud log response.

Sidewall core porosity values for the two intervals in the zone of interest range from 4.8 to 16.3 percent. However, permeability is extremely low, ranging from less than 0.1 to 3.0 md. The sidewall cores were analyzed, but no significant shows of hydrocarbons were reported. Both intervals within the zone of interest were tested on November 13, 1981. A drill stem test with perforations at 12,392 to 12,408 feet and at 12,418 to 12,430 feet recovered 564 barrels of water but no gas or liquid hydrocarbons (tables 8 and 9).

Depth (Feet)	Drilling Break	Sample Description (Mud Log)	Hot wire (units)	Chromatograph	Sidewall cores ^a	Well Log Interpretation ^b	Comments	Test
6,950	I	Siltstone with yellow mineral flu., White- yellow cut		1			,	1
11,000	I	Oolitic LS, and SS, dull yellow flu., dead oil?, no cut	I	ſ	I			ı
12,150	3 - 25	SS, firm, abundant brown stain, yellow flu., no cut	I	malfunction				T
12,400 -	9 - 16	LS, micrite, dull yellow - dark green flu., no cut	ı	с ₁	φ = 4.8-15.2% k =<0.1-1.8 md	$\phi = 23\%$ Sw = 25%	Characteristic crossover response of CNL / FDC,	DST of 12,392 - 12,408 and 12,418 - 12,430 ft recovered 564 bbl
14,428 - 14,433	9 - 25	LS, micrite, dull yellow - dark green flu., no cut	I	c ₁	φ = 6.1-16.3% k =<0.1-3.0 md	$\phi = 24\%$ Sw = 26\%	indicating presence of gas	of water

Table 9. Hydrocarbon show summary

a Sonic ϕ compares favorably with sidewall core ϕ . b Density ϕ used for Sw calculation.

GEOTHERMAL GRADIENT

Figure 8 shows bottomhole temperatures for five logging runs in the Exxon LC Block 133 No. 1 well plotted against depth. A temperature of 60 ^OF is assumed at the seafloor at an indicated depth of 310 feet (225-foot water depth plus 85-foot kelly bushing elevation). Shown also is a straight-line graph between the seafloor and total-depth temperatures in order to represent an overall geothermal gradient for the well, which is 1.18 degrees Fahrenheit per 100 feet. Calculated geothermal gradients for all Georges Bank wells range from 1.06 to 1.40 degrees Fahrenheit per 100 feet.



Figure 8. Well temperatures and geothermal gradient for the Exxon Lydonia Canyon Block 133 No. 1 well. Well temperatures from bottomhole temperatures of logging runs. Geothermal gradient based on bottomhole temperature of deepest logging run.

KEROGEN ANALYSIS

Taken and adapted from C. E. Fry, H. L. Cousminer, and J. K. Filer, MMS published report MMS 89-0007

Kerogen type and thermal rank were determined by microscopic examination of kerogen slides and palynology slides made from cuttings samples from the Exxon LC Block 133 No. 1 well. In this analysis, organic material is classified as one of four major types: algal-amorphous, organic material of marine origin, either recognizable algae or the unstructured remains of algal material; herbaceous, leafy portions of plants, including spores and pollen; woody, plant detritus with a lignified, ribbed structure; coaly, black opaque material, thought to be chemically inert. Visual estimates are made for the percentage of each type relative to the total abundance of kerogen contained in each of the slides. Algal material is generally considered the best source for oil; more structured terrestrial kerogen is primarily a gas source.

Thermal maturity of the organic material is estimated by comparing the color of various palynomorphs contained in the kerogen slides to the thermal alteration index (TAI) scale (figure 9) taken from Jones and Edison (1978). The colors displayed by the organic matter are an indication of the degree to which the kerogen has been thermally altered (Staplin, 1969).

Judging thermal maturity using samples from well cuttings must be done with great

care to ensure that the material being analyzed is indigenous to the level sampled. Caved or reworked materials will both give false indications of maturity. Oxidation caused by a highenergy environment of deposition also can alter the appearance of the organic material.

KEROGEN TYPE

Cuttings samples are available in the Exxon LC Block 133 No. 1 well below 640 feet, well depth below kelly bushing. Overall in the well, the abundance of coaly and woody kerogens is greater than that of herbaceous and algal kerogens, and algal material is much less abundant than the other three types (figure 10). To about 8,000 feet, the four kinds of kerogen are fairly constant in relative abundance, but below this depth, within the Middle Jurassic, coaly and woody types increase while herbaceous and algal types decrease in abundance. From about 13,000 feet to TD, 14,118 feet, woody and herbaceous abundance increases and coaly abundance decreases. However, within this interval, the herbaceous organic matter is degraded. Within Tertiary through Lower Jurassic rocks, above 8,000 feet, there are three intervals with significant algal kerogen abundance, 640 to 2,090 feet, 3,710 to 5,710 feet, and 6,970 to 7,600 feet. Within these intervals, two samples contained

Coal Rank	% Ro.	TAI	Spore Color	Principal Zones of Hydrocarbon Generation
Peat		1.0	Very Pale Yellow	
Lignite		- 2.0 - -	Pale Yellow Yellow	Immature
Sub-Bituminous C A C	0.5 -	- - - 2.5	Yellow- Orange	
B High Volatile Bituminous	10	- - -	Orange- Brown Reddish-	— — — — — — – – – – – – – – – – – – – –
Medium Volatile Bituminous	1.0 -	- 3.0 - - -	Brown Dark Reddish- Brown	Condensate and Wet Gas
Low Volative Bituminous	1.5 -	- 3.5 -	Dark Brown	
Semi - Anthracite Anthracite	2.0 - 2.5 - 3.0 - 3.5 - 4.0 -	- 3.7 - - 4.0	Black	Dry Gas

Figure 9. Relationships among coal rank, percent R_o, TAI, spore color, and thermal zones of hydrocarbon generation (after Jones and Edison, 1978).



Figure 10. Graph of kerogen types and organic thermal maturity for the Exxon Lydonia Canyon Block 133 No.1 well.

20 percent algal organic matter and most samples contained 10 percent or less.

THERMAL MATURITY

Thermal maturity estimates for hydrocarbon generation, based on palynomorph alteration colors, are presented in figure 10. A thermal alteration index (TAI) value of 2.7 occurs from 12,630 feet to total depth. This level of maturity is only marginally greater than the minimum threshold of 2.6. Shallower sedimentary units are immature for petroleum generation.

CONCLUSIONS

In the Exxon LC Block 133 No. 1 well, sedimentary rock above 12,630 feet is thermally immature for petroleum generation. Within the marginally mature interval below this depth, organic matter is mostly woody and coaly, although herbaceous kerogens increase to about 40 percent from 13,000 feet to total depth at 14,118 feet. However, organic richness is very poor. Figure 11 shows total organic carbon (TOC) analyses results for the well, and TOC below 12,600 feet is less than one percent for all samples. Altogether, the sedimentary section penetrated by this well does not appear to contain potential petroleum source rock.



Figure 11. Total organic carbon analysis for the Exxon Lydonia Canyon Block 133 No. 1 well. Data from Exxon.

PETROLEUM GEOCHEMISTRY

Taken and adapted from R. E. Miller, H. L. Cousminer, and C. E. Fry, MMS published report MMS 89-0007

The objective of this section is to assess the petroleum source potential of rocks penetrated by the Exxon LC Block 133 No. 1 well. Low-molecular-weight hydrocarbon analysis, total organic carbon, and thermal pyrolysis data are from Exxon Company, USA, Houston. Petrographic examination of organic residues and palynomorphs, in order to identify kerogen and extent of sedimentary thermal alteration, is reported in the **Kerogen analysis** chapter of this report.

Table 10. Total organic carbon

Depth (ft)	% TOC
4,770	0.88
6,970	0.57
7,270	0.93
7,570	0.82
7,870	1.12
8,170	1.43
8,270	1.24
8,370	1.30
8,770	1.33
8,770	1.35
9,070	1.31
9,070	1.51
9,270	1.46
9,470	1.52
9,570	1.42
9,670	1.19
9,870	0.77
10,070	0.57
10,370	1.01
10,570	1.02
10,770	0.98
10,970	0.47
11,170	0.60
11,370	1.52
11,470	0.60

ORGANIC RICHNESS

Source rock richness of the well is shown in figure 11 and table 10, which illustrate the distribution of total organic carbon (TOC) with depth for 50 well cuttings samples from 4,770 to 14,118 feet (TD).

Depth (ft)	% TOC
11,570	0.71
11,770	0.51
11,970	0.39
12,170	0.86
12,260	1.64
12,269	0.94
12,282	2.35
12,284	1.11
12,296	0.87
12,307	0.32
12,317	0.41
12,325	0.20
12,470	0.66
12,670	0.44
12,870	0.35
13,070	0.81
13,270	0.14
13,370	0.06
13,570	0.06
13,670	0.13
17,770	0.14
13,870	0.14
13,970	0.12
14,070	0.15
14,118	0.12

The Cretaceous section, from 1,180 to 6,520 feet, contains mixed lithologies but is not adequately represented with only one analysis, 0.88 weight percent TOC at 4,770 feet. The Jurassic section, from 6,520 to 13,100 feet, also consists of mixed lithologies and contains 40 TOC values that range from 0.35 to 2.35 and average 0.97 weight percent. Such values are in the poor to good range. Below 13,100 feet, in limestones that are undated but of probable Middle Jurassic age, TOC values decrease significantly and range from 0.06 to 0.14 weight percent. These values are poor. Thermal pyrolysis data, furnished by Exxon, show S₂ values less than one mg/g for all 50 cuttings samples. These values are especially low below 12,400 feet, less than 0.1 mg/g. Lowmolecular-weight hydrocarbon analyses, C_1 - C_4 , submitted by Exxon (figure 12, table 11) show low concentrations for the entire well (256 to 7,390 ppm). The highest concentrations are at shallow depths, above 2,000 feet, owing to biogenic methane. The gas concentrations generally decrease with depth in the lower part of the well, paralleling the total organic carbon and thermal pyrolysis results. Gasoline-range hydrocarbon concentrations are also low for the entire section penetrated by the well, ranging from zero to four ppm (table 11). The TOC, thermal pyrolysis, and lowmolecular-weight and gasoline-range hydrocarbon analyses all suggest that the entire sedimentary section is organically too lean to contain significant hydrocarbon source rocks.

KEROGEN TYPES

Types of organic matter present in the well are described in the **Kerogen analysis** chapter of this report. In three intervals the well contains up to 20 percent (of total kerogen) amorphous, algal-appearing material. These intervals are 640 to 2,090 feet, 3,710 to 5,710 feet, and 6,970-7,600 feet. The interval between 8,000 and 13,000 feet contains predominantly woody and coaly kerogens. From 13,000 feet to TD, herbaceous kerogens are in greater abundance, generally about 40 percent, although much of this organic matter has been degraded and is poorly preserved, losing most of its structure.

Thermal pyrolysis analysis by Exxon confirms that hydrogen-poor, woody and, coaly kerogens are the most abundant types of organic matter in this well. The hydrogen index, S₂/TOC, ranged from zero to 143. Values less than 200 suggest the dominance of gas-prone, terrestrially derived woody and coaly organic matter. The predominance of woody and coaly organic matter in this well is similar to the results reported for the COST G-1 and COST G-2 wells at equivalent depths (Miller and others, 1982). The Exxon LC Block 133 No. 1 well is located almost equidistant between the two COST wells.

THERMAL MATURITY

A thermal alteration index (TAI) value of 2.7 was recognized from 12,630 to 14,118 feet (TD) (figure 10). This level of thermal maturity is marginal for petroleum generation. Above 12,630 feet, the stratigraphic section is not thermally mature. Thermal pyrolysis transformation ratios $[S_1/S_1 + S_2)$] reported by Exxon do not show significant conversion of organic matter to petroleum in any samples for the entire depth of the well. However, with the low S_1 and S_2 values of these organically lean rocks, the transformation ratio is probably not a reliable indicator of thermal maturity, especially below 12,400 feet, where the section is extremely lean, but where the TAI data suggest that the



Figure 12. Low-molecular-weight hydrocarbon analysis for the Exxon Lydonia Canyon Block 133 No. 1 well. Data from Exxon.

Depth	C ₁ -C ₄	Gas Wetness	C ₄ -C ₇
(feet)	(ppm)	(%)	(ppm)
670	2,070.43	0.8527	
870	6,251.18	0.3823	
1,070	2,495.53	0.4853	
1,270	6,992.18	0.2845	
1,470	6,698.63	0.2974	
1,670	7,390.85	0.2818	
1,870	3,283.72	0.4611	
2,070	1,579.08	0.5579	
2,270	1,529.84	0.4634	
2,470	695.20	1.1004	
2,670	1,121.45	1.2573	
2,870	719.16	0.6786	
3,070	557.04	1.8311	
3,270	1,213.86	1.3156	
3,470	787.43	1.6725	
3,670	910.21	1.5799	
3,870	1,188.31	0.8264	
4,070	2,929.36	0.6554	
4,270	467.35	2.9378	
4,470	1,901.92	2.0474	
4,670	598.32	2.1794	
4,770	544.74	2.5609	0
4,870	351.36	2.3850	
4,970	523.50	3.8548	
5,070	461.19	3.3088	
5,170	1,202.03	2.0457	
5,270	918.57	4.3644	
5,370	1,283.84	3.9818	
5,470	552.92	4.5902	
5,570	901.19	3.4621	
5,670	255.69	2.4053	
5,770	1,276.27	3.9882	
5,870	973.02	4.3298	
5,970	2,052.23	2.5290	
6,070	2,646.16	4.2363	
6,170	4,845.61	5.7372	
6,270	4,645.07	3.6716	
6,370	3,357.97	7.8407	
6,470	3,052.37	4.8657	

Table 11. Low-molecular-weight and gasoline-range hydrocarbon analyses

continued

Depth	C ₁ -C ₄	Gas Wetness	C ₄ -C ₇
(feet)	(ppm)	(%)	(ppm)
6,570	1,524.71	8.2134	
6,670	2,472.42	4.4378	
6,770	874.45	6.4338	
6,870	995.21	12.6164	
6,970	684.26	15.1901	0
7,070	889.24	12.4972	
7,170	806.16	8.5306	
7,270	755.80	6.3416	0.976
7,370	1,274.02	4.4285	
7,470	782.79	11.5101	
7,570	811.53	10.2214	0
7,670	553.66	18.8376	
7,770	528.45	17.9411	
7,870	896.40	31.2082	0
7,970	886.72	27.7359	
8,070	920.66	15.1478	
8,170	2,601.72	6.9535	0
8,270	1,532.91	7.6580	1.077
8,370	1,403.54	5.6443	0
8,470	1,627.34	4.8742	
8,570	1,683.16	4.5688	
8,670	755.15	4.4044	1.232
8,770	442.14	12.9439	0.712
8,870	617.39	7.4442	
8,970	1,221.43	2.4185	
9,070	620.09	15.9993	2.397
9,170	1,034.40	4.3648	
9,270	872.92	3.4196	0.969
9,370	1,885.31	1.7371	
9,470	1,019.79	4.1979	0
9,570	2,641.56	1.9227	1.313
9,670	1,601.19	2.7086	0
9,770	1,875.15	1.5711	
9,870	2,183.09	2.2299	0
9,970	949.73	2.5776	
10,070	669.25	2.2144	0
10,170	1,451.46	1.6425	
10,270	833.60	4.2682	
10,370	942.44	3.5567	0
10,470	870.72	2.9401	
10,570	62.93	3.2791	0

 Table 11 Low-molecular-weight--continued

continued

Depth	C ₁ -C ₄	Gas Wetness	C ₄ -C ₇
(feet)	(ppm)	(%)	(ppm)
10,670	670.86	2.3343	
10,770	2,365.54	73.1867	0
10,870	1,666.65	5.8405	
10,970	765.61	7.2452	0
11,170	2,324.68	7.6140	0.473
11,270	719.86	14.6233	
11,370	650.72	25.2459	2.107
11,470	1,288.64	18.5646	1.229
11,570	1,743.35	8.3145	0.818
11,670	1,724.93	7.1586	
11,770	560.31	14.4099	0.867
11,870	538.23	17.9143	
11,970	503.87	8.2740	0
12,070	1,230.03	6.4096	
12,170	1,221.27	2.9166	0
12,270	1,379.46	3.6781	
12,370	563.38	6.8089	
12,470	303.45	5.6286	0
12,570	446.40	4.1465	
12,670	423.70	4.1444	0
12,770	391.51	7.9053	
12,870	615.13	3.8106	0
12,970	510.75	6.0891	
13,070	460.99	4.7572	1.982
13,170	562.26	4.4392	
13,270	360.64	12.2116	2.043
13,370	369.97	3.9273	0
13,570	405.08	3.2784	0
13,670	551.90	4.8070	0.882
13,770	600.75	3.8868	0
13,870	513.98	4.2881	0
13,970	551.27	4.3772	0
14,070	648.38	5.9070	4.482
14,118	1,447.69	2.4308	0

 Table 11 Low-molecular-weight--continued

section becomes marginally mature. Gas wetness (figure 12) does not increase at a particular depth and then remain elevated below that depth as an indication of thermally mature section. Rather, gas wetness values are variable throughout the well.

Among available data, depth to thermally mature sedimentary section is best indicated by the TAI analysis. With a marginally mature value of 2.7 at 12,630 feet, the bottom of the well appears to be in the upper part of the hydrocarbon evolution window. Peak generation would be deeper than the well.

Assuming 10 million years since onset of petroleum maturation, the timetemperature relationships proposed by Connan (1974) suggest that the temperature for the onset of petroleum generation is about 75 °C, and peak generation in the oil and gas zone is about 130 °C. With a geothermal gradient of 1.18 ^oF/100 ft (figure 8), the temperature for onset of generation is at a well depth of about 9,000 feet, and the depth of peak generation, about 17,000 to 18,000 feet. Among the assumptions of this approach is that the present-day geothermal gradient represents maximum subsurface temperatures through the Mesozoic and Cenozoic Eras. This is a poor assumption, especially in light of the intrusive igneous unit at 13,200 feet. However, if temperatures had been higher, the generation onset threshold would be shallower than 9,000 feet. The best available time/temperature data, from the TAI analysis, indicate immaturity above 12,630 feet. This apparent conflict

may suggest that onset of petroleum generation (given adequate source matter) was more recent than 10 million years.

BURIAL HISTORY

The burial history diagram for the stratigraphic section penetrated by the Exxon LC Block 133 No. 1 well (figure 13) is based on MMS biostratigraphy and on the Cretaceous and Jurassic time scales of Van Hinte (1976a and 1976b). In general, burial diagrams for Georges Bank wells show rapid Middle Jurassic subsidence followed by moderate subsidence in the Late Jurassic and Cretaceous and low burial rates through the Cenozoic. In those wells for which Neogene biostratigraphic data exist, indicated subsidence increases for this shallowest and most recent part of the section. The Exxon LC Block 133 No. 1 well data show rapid Jurassic and Cretaceous burial rates until the Maestrichtian age, and low rates for the rest of the Late Cretaceous and through the Tertiary until the Miocene age. The increased post-Miocene indicated burial rate is at least partly due to including water depth in the diagram.

In constructing figure 13, no adjustments have been made for sedimentary compaction or for section removed by erosion.

If only the deeper part of the identified Bathonian section is marginally mature for petroleum generation, it has become mature recently in geologic time, perhaps in less than the last ten million years.



Figure 13. Burial diagram for the Exxon Lydonia Canyon Block 133 No.1 well.

COMPANY-SUBMITTED DATA

Data and reports were submitted by Exxon Corporation, U.S.A., to MMS when the Exxon LC Block 133 No. 1 well was drilled, as required by Federal regulations and lease stipulations. Items of general geological, geophysical and engineering usefulness are listed below. Items not listed include routine submittals required by regulation, such as the Exploration Plan, Application for Permit to Drill, and daily drilling reports, and detailed operations information, such as drilling pressure and temperature data logs. Well "electric" logs are listed in the **Formation** Evaluation chapter. Listed and unlisted company reports and data are available through the Public Information Unit, Minerals Management Service, Gulf of

Mexico OCS Region, 1201 Elmwood Park Boulevard, New Orleans, Louisiana 70123-2394; telephone (504)736-2519 or 1-800-200-GULF, FAX (504)736-2620. Well logs are available on microfilm from the National Geophysical Data Center, 325 Broadway Street, Boulder CO 80303-3337, attn. Ms Robin Warnken; telephone (303)497-6338, FAX (303)497-6513; email rwarnken@NGDC.NOAA.GOV.

At a later date, additional original technical data, including well logs, will be added to the compact disk (CD) version of the Georges Bank well reports. The CD will be available from the Gulf of Mexico OCS Region Public Information Unit.

SELECTED COMPANY-SUBMITTED DATA

Physical formation (mud) log, Exploration Logging of USA, Inc., undated.

Seismic velocity survey (checkshot survey), Birdwell Division, Seismograph Service Corp., Tulsa OK, undated.

Velocity survey computation (well velocity and well seismic tool data), Schlumberger Ltd., Wireline Testing, Houston TX, undated. Core analysis report (sidewall cores), Erco Petroleum Services, Inc., Houston, September 14, 1981.

Conventional core description form (12,251-12,270 and 13,488 -13,491 feet), Exxon Co., U.S.A., Houston, September 18, 1981.

Geochemical analyses (C_1 - C_4 and C_4 - C_7 hydrocarbon, total organic carbon, and thermal pyrolysis analyses), Exxon Co., U.S.A., Houston, undated.

SELECTED REFERENCES

This list is compiled from published and unpublished Minerals Management Service and USGS Conservation Division reports on Georges Bank wells. Not all of the references could be located and verified.

- Albrecht, P., 1970, Etude de constituents organiques des series sedimentaries de Logbaba et Messel. Transformations deagenetiques: Universite de Strasbourg, Memoires du Service de la Carge Geologique d'Alsac et de Lorraine, no. 32, 119 p.
- Amato, R.V. and J.W. Bebout, 1978, Geological and Operational Summary, COST No. GE-1 Well, Southeast Georgia Embayment Area, South Atlantic OCS: U. S. Geological Survey Open-File Report 78-668, 122 p.
- Amato, R. V. and J. W. Bebout (eds.), 1980, Geologic and Operational Summary, COST No. G-1 Well, Georges Bank Area, North Atlantic OCS: U. S. Geological Survey Open-File Report 80-268, 112 p.
- Amato, R.V., and E.K. Simonis (eds.), 1979, Geologic and Operational Summary, COST No. B-3 Well, Baltimore Canyon Trough Area, Mid-Atlantic OCS: U.S. Geological Survey Open-File Report 79-1159, 118 p.
- Amato, R.V. and E.K. Simonis,(eds.), 1980, Geologic and Operational Summary, COST No. G-2 Well, Georges Bank Area, North Atlantic OCS: U. S. Geological Survey Open-File Report 80-269, 116 p.
- BBN-Geomarine Services Co., 1975, COST wellsite G-1, Georges Bank, engineering geology interpretation of high-resolution geophysical data: Houston, Texas, 11 p.
- Ballard, R. D. and E. Uchupi, 1975, Triassic rift structure in Gulf of Maine: American Association of Petroleum Geologists Bulletin, v. 59, no. 7, p. 1041-1072.
- Bayliss, G. S., 1980, Source-rock evaluation reference manual: Houston, Texas, Geochem Laboratories, Inc., 80 p.
- Bebout, J. W., 1980, Observed stratigraphic distribution of spores, pollen, and *incertae sedis* palynomorphs in the Tertiary section of the COST No. B-2 well, Baltimore Canyon, Atlantic Outer Continental Shelf: Palynology, v. 4, p. 181-196.
- Bebout, J. W., 1981, An informal palynologic zonation for the Cretaceous System of the United States Mid-Atlantic (Baltimore Canyon area) Outer Continental Shelf: Palynology, v. 5, p. 159-194.
- Berggren, W.A., D.V. Kent, C.C. Swisher III, and M.P. Aubry, 1995, A revised Cenozoic geochronology and chronostratigraphy; *in* Geochronology Time Scales and Global Stratigraphic Correlation, SEPM Special Publication no. 54, p. 129-212.
- Bhat, H., N. J. McMillan, J. Aubert, B. Porthault, and M. Surin, 1975, North American and African drift--the record in Mesozoic coastal plain rocks, Nova Scotia and Morocco, *in* Yorath, C. J., E. R. Parker, and D. J. Glass, (eds.), Canada's Continental Margins and Offshore Petroleum Exploration: Canadian Society of Petroleum Geologists Memoir 4, p. 375-389.
- Brideau, W. W. and W. C. Elsick, (eds.), 1979, Contributions of stratigraphic palynology (v. 2), Mesozoic Palynology: American Association of Stratigraphic Palynologists Contributions Series No. 4.

- Bronnimann, P., 1955, Microfossils *incertae sedis* from the Upper Jurassic and Lower Cretaceous of Cuba: Micropaleontology, v. 1, pp. 28, 2 pl., 10 text.
- Bujak, J. P., M. S. Barss, and G. L. Williams, 1977, Offshore east Canada's organic type and color and hydrocarbon potential: Oil and Gas Journal, v. 75, no. 15, p. 96-100.
- Bujak, J. P. and M. J. Fisher, 1976, Dinoflagellate cysts from the Upper Triassic of Arctic Canada: Micropaleontology, v. 22, p. 44-70, 9 pls.
- Bujak, J. P, and G. L. Williams, 1977, Jurassic palynostratigraphy of offshore eastern Canada, *in* Swain, F. M., (ed.), Stratigraphic Micropaleontology of Atlantic Basin and Borderlands: New York, Elsevier Scientific Publishing Co., p. 321-339.
- Bukry, D., 1969, Upper Cretaceous coccoliths from Texas and Europe: University of Kansas Paleontological Contributions, Art. 5 (Protista 2), p. 1-9, 50 pl., 1 text.
- Burk, C. A. and C. L. Drake, (eds.), 1974, Geology of Continental Margins: New York, Springer-Verlag, 1,009 p.
- Burke, K., 1975, Atlantic evaporites formed by evaporation of water spilled from Pacific, Tethyan, and southern oceans: Geology, v. 3, no. 11, p. 613-616.
- Cepek, P. and W. W. Hay, 1970, Zonation of the Upper Cretaceous using calcareous nannoplankton: Palaontologische Abhandlungen, Abtelung B Palabotanik, Band III, Heft 3/4, p. 333-340.
- Cita, M. B. and S. Gartner, 1971, Deep Sea Upper Cretaceous from the western North Atlantic: *in* Proceedings II International Planktonic Conference, Roma, 1970: Rome, Edizioni Tecnoscienza, v. 1, p. 287-319.
- Clarke, R. F. A, and J. P. Verdier, 1967, An investigation of microplankton assemblages from the chalk of the Isle of Wight, England: Verhandelingen der Koninklijke Nederlandische Akademie van Wetenschappen, Afdeeling Natuurkunde, and Eerste Reeks, 24, p. 1-96.
- Claypool, G. E., C. M. Lubeck, J. P. Baysinger, and T. G Ging, 1977, Organic geochemistry, *in* Scholle, P. A., (ed.), Geological studies on the COST No. B-2 well, U. S. Mid-Atlantic Outer Continental Shelf area: U. S. Geological Survey Circular 750, p. 46-59.
- Connan, J. 1974, Time-temperature relation in oil genesis: American Association of Petroleum Geologists Bulletin, v. 58, no. 12, p. 2516-2521.
- Core Laboratories, Inc., 1976, Core studies, COST Atlantic well No. G-1, Georges Bank, Offshore Atlantic Ocean: Dallas, Texas, 153 p.
- Core Laboratories, Inc., 1977a, Core studies, COST Atlantic well No. G-2, Georges Bank, Offshore Atlantic Ocean: Dallas, Texas, 298 p.
- Core Laboratories, Inc., 1977b, Geochemical service report, COST G-2 Atlantic well, Georges Bank, offshore Massachusetts, U. S. A.: Dallas, Texas, 147 p.
- Council on Environmental Quality, 1974, OCS oil and gas--An environmental assessment--A report to the President by the Council on Environmental Quality: Washington, D. C. (U. S. Government Printing Office), Stock No. 4000-00322, v. 1, 214 p.

- Cousminer, H. L., 1984, Canadian dinoflagellate zones (Middle Jurassic to Middle Eocene) in Georges Bank Basin (abstract): Proceedings of the American Association of Stratigraphic Palynologists, Arlington, Virginia, v. 9, p. 238.
- Cousminer, H. L., W. E. Steinkraus, and C. E. Fry, 1982, Biostratigraphy and thermal maturation profile, Exxon 133 No. 1 (OCS-A-0170) well section: Unpublished Report, Minerals Management Service.
- Cousminer, H. L., W. E. Steinkraus, and R. E. Hall, 1984, Biostratigraphic restudy documents Triassic/Jurassic section in Georges Bank COST G-2 well (abstract): Proceedings of the American Association of Petroleum Geologists, Annual Meeting, San Antonio, Texas, v. 68, no. 4, p. 466.
- Davey, R. J., 1979, The stratigraphic distribution of dinocysts in the Portlandian (latest Jurassic) to Barremian (Early Cretaceous) of northwest Europe: American Association of Stratigraphic Palynologists Contributions, Series No. 5B, p. 49-81.
- Davey, R. J. and J. P. Verdier, 1974, Dinoflagellate cysts from the Aptian type sections at Gargas and La Bedoule, France: Paleontology, v. 17, pt. 3, p. 623-653.
- Davies, E. H., 1985, The miospore and dinoflagellate cyst oppel-zonation of the Lias of Portugal: Palynology, v. 9, p. 105-132.
- Dorhofer, G. and E. H. Davies, 1980, Evolution of archeopyle and tabulation in Rhaetogonyaulacian dinoflagellate cysts: Royal Ontario Museum, Life Sciences Miscellaneous Publications, p. 1-91, fig. 1-40.
- Dow, W. G., 1974, Application of oil-correlation and source-rock data to exploration in Williston Basin: American Association of Petroleum Geologists Bulletin, v. 58, no. 7, p. 1253-1262.
- Dow, W. G., 1977, Kerogen studies and geological interpretations: Journal of Geochemical Exploration, v. 7, p. 79-99.
- Drake, C. L., J. I. Ewing, and H. Stockard, 1968, The continental margin of the eastern United States: Canadian Journal of Earth Science, v. 5, no. 4, p. 993-1010.
- Drake, C. L., M. Ewing, and G. H. Sutton, 1959, Continental margins and geosynclines--The east coast of North America north of Cape Hatteras, *in* Aherns, L. H., and others, (eds.), Physics and Chemistry of the Earth, v. 3: New York, Pergamon, p. 110-198.
- Eliuk, L. S., 1978, the Abenaki Formation, Nova Scotia, Canada--A depositional and diagenetic model for a Mesozoic carbonate platform: Bulletin of Canadian Petroleum Geology, v. 26, no. 4, p. 424-514.
- Emery, K. O. and E. Uchipi, 1972, Western North Atlantic Ocean--Topography, rocks, structure, water, life, and sediments: American Association of Petroleum Geologists Memoir 17, 532 p.
- Evitt, W. R., (ed.), 1975, Proceedings of a forum on dinoflagellates: American Association of Stratigraphic Palynologists Contributions, Series No. 4, 76 p.
- Folger, D. W., 1978, Geologic hazards on Georges Bank--an overview: Geological Society of America Abstracts with Programs, v. 10, no. 1, p. 42.
- Fry, C. E., 1979, Geothermal gradient, *in* Amato, R. V. and E. K. Simonis (eds.), Geologic and Operational Summary, COST No. B-3 well, Baltimore Canyon Trough Area, Mid-Atlantic OCS: U. S. Geological Survey Open-File Report 79-1159, p. 64-65.

- Gartner, S., Jr., 1968, Coccoliths and related calcareous nannofossils from Upper Cretaceous deposits of Texas and Arkansas: University of Kansas Paleontological Contributions, no. 48, Protista, v. 48, Art. 1, p. 1-56.
- GeoChem Laboratories, Inc., 1976, Hydrocarbon source facies analysis, COST Atlantic G-1 well, Georges Bank, offshore Eastern United States: Houston, Texas, 10 p.
- GeoChem Laboratories, Inc., 1977, Hydrocarbon source facies analysis, COST Atlantic G-2 well, Georges Bank, offshore eastern United States: Houston, Texas, 66 p.
- Gibson, T. G., 1970, Late Mesozoic-Cenozoic tectonic aspects of the Atlantic coastal margin: Geological Society of America Bulletin, v. 81, no. 6, p. 1813-1822.
- Gitmez, G. U. and W. A. S. Sarjeant, 1972, Dinoflagellate cysts and acritarchs from the Kimmeridgian (Upper Jurassic) of England, Scotland and France: Bulletin of the British Museum of Natural History: Geology, v. 21, p. 171-257.
- Given, M. M., 1977, Mesozoic and Early Cenozoic geology of offshore Nova Scotia: Bulletin of Canadian Petroleum Geology, v. 25, p. 63-91.
- Gocht, H., 1970, Dinoflagellaten-Zysten aus dem Bathonium des erdolfeldes Aldorf (Northwest-Setuschland): Palaeontographics, Abt. B., v. 129, p. 125-165.
- Gorka, H., 1963, Coccolithophorides, Dinoflagellates, Hystrichosphaerides et microfossiles *incertae sedis* du Cretace superier de Pologne: Acta Palaeontological Polonica, v. 8, p. 1-82.
- Gradstein, F.M., F.P.Achterberg, J.G. Ogg, J.Hardenbol, P. van Veen, and Z. Huang, 1995, A Triassic, Jurassic, and Cretaceous time scale; *in* Geochronology Time Scales and Stratigraphic Correlation, SEPM Special Publication no. 54, p. 95-126.
- Grose, P. L. and J. S. Mattson, 1977, The Argo Merchant oil spill--A preliminary scientific report: National Oceanic and Atmospheric Administration Environmental Research Laboratories, 129 p.
- Grow, J. A., R. E. Mattick, and J. S. Schlee, 1979, Multichannel seismic depth sections and interval velocities over continental shelf and upper continental slope between Cape Hatteras and Cape Cod, *in* Watkins, J. S., L. Montadert, and P. W. Dickerson, (eds.), Geological and Geophysical Investigations of Continental Margins: American Association of Petroleum Geologists Memoir 29, p. 65-83.
- Harwood, R. J., 1977, Oil and gas generation by laboratory pyrolysis of kerogen: American Association of Petroleum Geologists Bulletin, v. 61, no. 12, p. 2082-2102.
- Hill, M. E., III, 1976, Lower Cretaceous Nannofossils from Texas and Oklahoma: Paleontographica, Abtelung B, 156, Lfg. 4-6, p. 103-179.
- Hunt, J. M., 1967, The origin of petroleum in carbonate rocks: *in* G. V. Chilingar, H. S. Bissell, and R. W. Fairbridge, (eds.), Carbonate Rocks: New York, Elsevier, p. 225-251.
- Hunt, J. M., 1974, Hydrocarbon and kerogen studies, in C. C von der Borch and others, Initial Reports of the Deep Sea Drilling Project, v. 22: Washington, D. C., U. S. Government Printing Office, p. 673-675.
- Hunt, J. M., 1978, Characterization of bitumens and coals: American Association of Petroleum Geologists Bulletin, v. 62, no. 2, p. 301-303.
- Hunt, J. M., 1979, Petroleum Geochemistry and Geology: San Francisco, W. H. Freeman Co., p. 273-350.

- Hurtubise, D. O. and J. H. Puffer, 1985, Nepheline normative alkalic dolerite of the Georges Bank Basin, North Atlantic, part of an Early Cretaceous eastern North American alkalic province: Geological Society of America, Northeastern Section, 20th Annual Meeting, 1985, v. 17, no. 1, p. 25.
- Hurtubise, D. O., J. H. Puffer, and H. L. Cousminer, 1987, An offshore Mesozoic igneous sequence, Georges Bank Basin, North Atlantic: Geological Society of America Bulletin, v. 98, no. 4, p. 430-438.
- International Biostratigraphers, Inc., 1976, Biostratigraphy of the COST G-1 Georges Bank test: Houston, Texas, 16 p.
- International Biostratigraphers, Inc., 1977, Biostratigraphy of the COST G-2 Georges Bank test: Houston, Texas, 16 p.
- Jansa, L. F. and J. A. Wade, 1975, Geology of the continental margin off Nova Scotia and Newfoundland, *in* W. J. M van der Linden and J. A. Wade (eds.), Offshore Geology of Eastern Canada: Geological Survey of Canada Paper 74-30, v. 2, p. 51-105.
- Jansa, L. F. and J. Wiedmann, 1982, Mesozoic-Cenozoic development of the eastern North American and northwest African continental margins: a comparison, *in* V. von Rad, K. Hinz, M.Sarnthein, and E. Seibold (eds.), Geology of the Northwest African Continental Margin: Berlin, Springer-Verlag, p. 215-269.
- Jansa, L. F., G. L. Williams, J. A. Wade, and J P. Bujak, 1978, COST B-2 well (Baltimore Canyon) and its relation to Scotian Basin (abstract): American Association of Petroleum Geologists Bulletin, v. 62, no. 3, p. 526.
- Jones, R. W. and T. A. Edison, 1978, Microscopic observations of kerogen related to geochemical parameters with emphasis on thermal maturation, *in* D. F Oltz (ed.), Geochemistry: Low Temperature Metamorphism of Kerogen and Clay Minerals: Society of Economic Paleontologists and Mineralogists, Pacific Section, Annual Meeting, Los Angeles, p. 1-12.
- Kent, D. V. and F. M. Gradstein, 1986, A Jurassic to Recent chronology, *in* P. R. Vogt and B. E. Tucholke (eds.), The Geology of North America, vol. M, The Western North Atlantic Region: Geological Society of America, p. 45-50.
- King, L. H. and B. MacLean, 1975, Geology of the Scotian Shelf and adjacent areas: Canadian Geological Survey Paper 74-23, p. 22-53.
- Kinsman, D. J. J., 1975, Rift Valley basins and sedimentary history of trailing continental margins, in A. G. Fisher and S. Judson, (eds.), Petroleum and Global Tectonics: Princeton, Princeton University Press, p. 83-126.
- Kjellstrom, G., 1973, Maastrichtian microplankton from the Hollviken borehole No. 1 in Scania, southern Sweden: Sveriges Geologiska Undersokning, Afhandligar och Uppsatser, v. 7, p. 1-59.
- Landes, K. K. 1967, Eometamorphism and oil and gas in time and space: American Association of Petroleum Geologists Bulletin, v. 51, no. 6, p. 828-841.
- LaPlante, R. E., 1974, Hydrocarbon generation in Gulf Coast tertiary sediments: American Association of Petroleum Geologists Bulletin, v. 58, no. 7, p. 1281-1289.
- Larskaga, Ye. S. and D. V. Zhabreu, 1964, Effects of stratal temperatures and pressures on the composition of dispersed organic matter (from the example of the Mesozoic-Cenozoic deposits of the Western Ciccaspian region): Dokl. Akad. Nauk SSSR, v. 157, no. 4, pp. 135-139.

- Lentin, J. K. and G. L. Williams, 1981, Fossil Dinoflagellates, Index to Genera and Species: Bedford Institute of Oceanography Report Series B1-R-81-12, p. 1-345.
- Louis, M. C. and B. P. Tissot, 1967, Influence de la temperature et de la pression sur la formation des hydrocarbures dans les argiles a kerogen [Influence of temperature and pressure on the generation of hydrocarbons in shales containing kerogen], *in* 7th World Petroleum Congress, Proceedings, (Mexico), v. 2: Chichester, International, John Wiley and Sons, p. 47-60.
- Lowell, J. D., G. J. Genik, T. H. Nelson, and P. M. Tucker, 1975, Petroleum and plate tectonics of the southern Red Sea, *in* A. G Fisher and S. Judson, (eds.), Petroleum and Global Tectonics: Princeton University Press, Princeton, p. 129-153.
- McIver, N. L., 1972, Cenozoic and Mesozoic stratigraphy of the Nova Scotia shelf: Canadian Journal of Earth Sciences, v. 9, p. 54-70.
- MacLean, B.C., and J.A. Wade, 1992, Petroleum geology of the continental margin south of the islands of St. Pierre and Miquelon, offshore eastern Canada; Bulletin of Canadian Petroleum Geology, v. 40, no. 3, p. 222-253.
- Maher, J. C., 1971, Geologic Framework and Petroleum Potential of the Atlantic Coastal Plain and Continental Shelf: U. S. Geological Survey Professional Paper 659, 98 p.
- Martini, E., 1971, Standard Tertiary and Quaternary calcareous nannoplankton zonation *in* Proceedings II International Planktonic Conference, Roma, 1970: Rome, Edizioni Tecnoscienza, p. 739-785.
- Mattick, R. E., R. Q. Foote, N. L. Weaver, and M. S. Grim, 1974, Structural framework of United States Atlantic Outer Continental Shelf north of Cape Hatteras: American Association of Petroleum Geologists Bulletin, v. 58, no. 6, 1179-1190.
- Miller, R. E., H. E. Lerch, G. E. Claypool, M. A. Smith, D. K. Owings, D. T. Lignon, and S. B. Eisner, 1982, Organic geochemistry of the Georges Bank basin COST Nos. G-1 and G-2 wells, *in* P. A. Scholle and C. R. Wenkam (eds.), Geological Studies of the COST Nos. G-1 and G-2 Wells, Unites States North Atlantic Outer Continental Shelf: U. S. Geological Survey Circular 861, p. 105-142.
- Miller, R. E., R. E. Mattick, and H. E. Lerch, 1981, Petroleum geochemistry and geology of Cenozoic and Mesozoic sedimentary rocks from Georges Bank basin (abstract): American Association of Petroleum Geologists Bulletin, v. 65, no. 9, p. 1667.
- Miller, R. E., D. M. Schultz, G. E. Claypool, H. E. Lerch, D. T. Lignon, C. Gary, and D. K. Owings, 1979, Organic geochemistry, *in*, P. A Scholle (ed.), Geological Studies of the COST GE-1 Well, United States South Atlantic Outer Continental Shelf Area: U. S. Geological Survey Circular 800, p. 74-92.
- Miller, R. E., D. M. Schultz, G. E. Claypool, M. A. Smith, H. E. Lerch, D. Ligon, D. K. Owings, and C. Gary, 1980, Organic geochemistry, *in* P.A. Scholle (ed.), Geological Studies of the COST No. B-3 Well, United States Mid-Atlantic Continental Slope Area: U. S. Geological Survey Circular 833, p. 85-104.
- Miller, R. E., D. M. Schultz, H. E. Lerch, D. T. Lignon, and P. C. Bowker, 1986, *in* Edson, G. M.(ed.), Shell Wilmington Canyon 586-1 Well, Geological and Operation Summary: Minerals Management Service, OCS Report MMS 86-0099, p. 37-44.
- Miller, R. E., D. M. Schultz, H. E. Lerch, D. T. Lignon, and P. C. Bowker, 1987, *in* Edson, G. M. (ed.), Shell Wilmington Canyon 587-1 Well, Geological and Operation Summary: Minerals Management Service, OCS Report MMS 87-0074, p. 39-46.

- Momper, J. A., 1978, Oil migration limitations suggested by geological and geochemical considerations, *in* Physical and Chemical Constraints on Petroleum Migration: American Association of Petroleum Geologists, Continuing Education Course Note Series No., 8, p. B1-B60.
- Morbey, S. J., 1975, The palynostratigraphy of the Rhaetian Stage Upper Triassic in the Kerdelbachgraben Austria: Paleontographica Abtrlung B, v. 152, p. 1-75, p. 1-19.
- Murray, G. E., 1961, Geology of the Atlantic and Gulf Coastal Provinces of North America: New York, Harper, 692 p.
- Orr, W. L., 1974, Changes in sulfur content and isotopic ratios of sulfur during petroleum maturation--study of Big Horn Basin Paleozoic oils: American Association Petroleum Geologists Bulletin, v. 58, no. 11, p. 2295-2318.
- Perry, W. J., J. P. Minard, E. G. A. Weed, E. I. Robbins, and E. C. Rhodehamel, 1975, Stratigraphy of the Atlantic continental margin of the United States north of Cape Hatteras--brief survey: American Association of Petroleum Geologists Bulletin, v. 59, no. 9, p. 1529-1548.
- Phillipi, G. T., 1957, Identification of oil-source beds by chemical means, *in* 20th International Geological Congress Proceedings: Mexico City (1956), Sec. 3, p. 25-38.
- Phillipi, G.T., 1965, On the depth, time, and mechanism of petroleum generation: Geochim. Cosmochim. Acta, v. 29, p. 1021.
- Postuma, J. A., 1971, Manual of Planktonic Foraminifera: New York, Elsevier, 420 p.
- Pusey, W. C., III, 1973, The ESR-kerogen method--how to evaluate potential gas and oil source rocks: World Oil, v. 176, no. 5, p. 71-75.
- Reinhardt, P., 1966, Zur taxonomie and biostratigraphie des fossilen nannoplanktons aus dem Malm, der Kreide und dem Alttertiar Mitteleuropas [Taxonomy and biostratigraphy of Malm, Cretaceous, and early Tertiary nannoplanktoanic faunas of central Europe], Frieberger Forschungshefte, Reihe C: Geowissenschaften, Mineralogie-Geochemie, 196 Paleont.: Leipzig, Bergakademie Freiberg, p. 5-61.
- Ricciardi, K. (ed.), 1989, Exxon Lydonia Canyon 133-1 Well, Geological and Operational Summary: Minerals Management Service OCS Report MMS 89-0007, 46 p.
- Riding, J. B., 1984, Dinoflagellate cyst range-top biostratigraphy of the uppermost Triassic to lowermost Cretaceous of northwest Europe: Palynology, v.8, p. 195-210.
- Robbins, E. I. and E. C. Rhodehamel, 1976, Geothermal gradients help predict petroleum potential of Scotian Shelf: Oil & Gas Journal, v. 74, no. 9, p. 143-145.
- Rona, P. A., 1973, Relations between rates of sediment accumulation on continental shelf, sea-floor spreading, and eustasy inferred from central North Atlantic: Geological Society of America Bulletin, v. 84, no. 9, p. 2851-2872.
- Ryan, W. B. F., M. B. Cita, R. L. Miller, D. Hanselman, B. Hecker, and M. Nibbelink, 1978, Bedrock geology in New England submarine canyons: Oceanologia Acta, v. 1, no. 2, p. 233-254.
- Sarjeant, W. A. S., 1979, Middle and Upper Jurassic dinoflagllate cysts--the world excluding North America: American Association of Stratigraphic Palynologists Contributions Series no. 5-B, p. 133-157.

- Schlee, J. S., J. C. Behrendt, J. A. Grow, J. M. Robb, R. E. Mattick, P. T. Taylor, and B. J. Lawson, 1976, Regional geologic framework off northeastern United States: American Association of Petroleum Geologists Bulletin, v. 60, no. 6, p. 926-951.
- Schlee, J. S., W. P. Dillon, and J. A. Dillon, 1979, Structure of the continental slope off the eastern United States, *in* L. J.Doyle and O. H. Pilkey, (eds.), Geology of Continental Slopes: Society of Economic Paleontologists and Mineralogists Special Publication 27, p. 95-117.
- Schlee, J.S. and K.D. Klitgord, 1988, Georges Bank basin: a regional synthesis; in R.E. Sheridan and J.A. Grow (eds.), The Geology of North America, vol. I-2, The Atlantic Continental Margin, Geological Society of America, p. 243-268.
- Schlee, J. S., R. G. Martin, R. E. Mattick, W. P. Dillon, and M. M. Ball, 1977, Petroleum geology of the U. S. Atlantic--Gulf of Mexico margins, *in* V. S Cameron (ed.), Exploration and Economics of the Petroleum Industry--New Ideas, Methods, New Developments: Southwestern Legal Foundation: New York, Mathew Bender and Co., v. 15, p. 47-93.
- Schlee, J. S., R. E. Mattick, D. J. Taylor, O. W. Girard, E. C., Rhodehamel, W. J. Perry, and K. C. Bayer, 1975, Sediments, structural framework, petroleum potential, environmental conditions and operation considerations of the United States North Atlantic Outer Continental Shelf: U. S. Geological Survey, Open-File Report 75-353, 179 p.
- Scholle, P. A. and C. R. Wenkam (eds.), 1982, Geological studies of the COST Nos. G-1 and G-2 wells, United States North Atlantic OCS: U. S. Geological Survey Circular 861, 193 p.
- Schultz, L. K. and R. L. Grover, 1974, Geology of Georges Bank Basin: American Association of Petroleum Geologists Bulletin, v. 58, no. 6, p. 1159-1168.
- Schwab, K.W., P. van Gijzel, and M.A. Smith, 1990, Kerogen evolution and microscopy workshop short course, International Symposium on Organic Petrology, Zeist, the Netherlands, January 10 and 11, 1990 (unpublished).
- Shell Canada Limited, 1970a, Well history report, Oneida O-25, 50 p.
- Shell Canada Limited, 1970b, Well history report, Mohawk B-93, 25 p.
- Shell Canada Limited, 1972, Well history report, Mohican I-100, 76 p.
- Sheridan, R. E., 1974a, Conceptual model for the block-fault origin of the North American Atlantic continental margin geosyncline: Geology, v. 2, no. 9, p. 465-468.
- Sheridan, R. E., 1974b, Atlantic continental margin of North America, *in* C. A. Burk and C. L. Drake, (eds.), Geology of Continental Margins: New York, Springer-Verlag, p. 391-407.
- Sheridan, R. E., 1976, Sedimentary basins of the Atlantic margin of North America: Tectonophysics, v. 36, p. 113-132.
- Sherwin, D. F., 1973, Scotian Shelf and Grand Banks, in R. G. McCrossan (ed.), Future Petroleum Provinces of Canada--Their Geology and Potential: Canadian Society of Petroleum Geologists Memoir 1, p. 519-559.
- Singh, C., 1971, Lower Cretaceous microfloras of the Peace River area, northwestern Alberta: Research Council of Alberta Bulletin 28, 2 volumes, 542 p.

- Smith, H. A., 1975, Geology of the West Sable structure: Bulletin of Canadian Petroleum Geology, v. 23, no. 1, p. 109-130.
- Smith, M. A., 1979, Geochemical analysis, *in* R. V. Amato and E. K. Simonis (eds.), Geologic and Operational Summary, COST No. B-3 Well, Baltimore Canyon Trough Area, Mid-Atlantic OCS: U. S. Geological Survey Open-File Report 79-1159, p. 81-99.
- Smith, M. A., 1980, Geochemical analysis, *in* R.V. Amato and E.K. Simonis (eds.), Geologic and Operational Summary, COST No. G-2 Well, Georges Bank Area, North Atlantic OCS: U. S. Geological Survey Open-File Report, 80-269, p. 77-99.
- Smith, M. A., 1995, Assessment of U.S. Atlantic hydrocarbon resources using new geochemical technology: U.S. Geological Society of America, Abstracts with programs, 1995 Annual Meeting, New Orleans, LA.
- Smith, M.A., R.V. Amato, M.A. Furbush, D.M. Pert, M.E. Nelson, J. S. Hendrix, L.C. Tamm, G. Wood, Jr., and D.R. Shaw, 1976, Geological and Operational Summary, COST No. B-2 Well, Baltimore Canyon Trough Area, Mid-Atlantic OCS: U. S. Geological Survey Open-File Report 76-774, 79 p.
- Smith, M. A. and D. R. Shaw, 1980, Geochemical analysis, *in* R. V. Amato and J. W. Bebout (eds.), Geologic and Operational Summary, COST No. G-1 well, Georges Bank Area, North Atlantic OCS: U. S. Geological Survey Open-File Report 80-268, p. 81-94.
- Smith, M.A., and P. van Gijzel, 1990, New perspectives on the depositional and thermal history of Georges Bank; *in* W.J.J. Fermont and J.W. Weegink (eds.), Proceedings, International Symposium on Organic Petrology, Zeist, the Netherlands.
- Smith, R. A., J. R. Stack, and R. K. Davis, 1976, An oil spill risk analysis for the Mid-Atlantic Outer Continental Shelf lease area: U. S. Geological Survey Open-File Report 76-451, 24 p.
- Staplin, F. L., 1969, Sedimentary organic matter, organic metamorphism, and oil and gas occurrence: Bulletin of Canadian Petroleum Geology, v. 17, no. 1, p. 47-66.
- Steinkraus, W. E., 1980, Biostratigraphy, in R. V. Amato and J. W. Bebout, (eds.), Geologic and Operation Summary, COST No. G-1 Well, Georges Bank, North Atlantic OCS: U. S. Geological Survey Open-File Report 80-268, p. 39-51.
- Stewart, H. B., Jr. and G. F. Jordan, 1964, Underwater sand ridges on Georges Shoal, *in* R. L. Miller (ed.), Papers in Marine Geology, Shepard Commemorative Volume: New York, Macmillan, p. 102-114.
- Tamm, L. C., 1978, Electric log interpretations, *in* R. V. Amato and J. W. Bebout (eds.), Geological and Operational Summary, COST No. GE-1 Well, Southeast Georgia Embayment Area, South Atlantic OCS: U. S. Geological Survey Open-File Report 78-668, 61-75.
- Thierstein, H. R., 1971, Tentative Lower Cretaceous calcareous nannoplankton zonation: Ecolgae Geologicae Helvetiae, v. 64, p. 459-487.
- Tissolt, B. P. and D. H. Welte, 1978, Petroleum Formation and Occurrence, A New Approach to Oil and Gas Exploration: Berlin, Springer-Verlag, p. 123-201.
- Tissot, B., B. Durand, J. Espitalie, and A. Combaz, 1974, Influence of nature and digenesis of organic matter in formation of petroleum: American Association of Petroleum Geologists Bulletin, v. 58, no. 3, p. 499-506.

- Tschudy, R. H., 1973, *Complexiopollis* Pollen Lineage in Mississippi Embayment Rocks: U. S. Geological Survey Professional Paper 743-C, p. C1-C15.
- Uchupi E. and K. O. Emery, 1967, Structure of continental margin off Atlantic coast of United States: American Association of Petroleum Geologists Bulletin, v. 51, no. 2, p. 223-234.
- U. S. Department of Commerce, 1973, Environmental Conditions within Specified Geographical Regions--Offshore East and West Coast of the United States and in the Gulf of Mexico: Washington, D. C., National Oceanographic Data Center, National Oceanographic and Atmospheric Administration, 735 p.
- Van Gijzel, P., 1990, Transmittance colour index (TCI) of amorphous organic matter: a new thermal maturity indicator for hydrocarbon source rocks, and its correlation with mean vitrinite reflectance and thermal alteration index (TAI); *in* W.J.J. Fermont and J.W. Weegink, eds., Proceedings, International Symposium on Organic Petrology, Zeist, the Netherlands.
- Van Hinte, J. E., 1976a, A Jurassic time scale: American Association of Petroleum Geologists Bulletin, v. 60, no. 4, p. 489-497.
- Van Hinte, J. E., 1976b, A Cretaceous time scale: American Association of Petroleum Geologists Bulletin, v. 60, no. 4, p. 498-516.
- Vassoyevich, N. B., Yu. I. Korchagina, N. V. Lopatin, and V. V. Chernyshev, 1969, Glavanaya faza nefteobrazovaniya [Principal phase of oil formation]: Moskovskogo Universiteta Vestnik, Ser. 4, Geologii, v. 24, no. 6, p. 3-27: English translation *in* International Geology Review, 1970, v. 12, no. 11, p. 1,276-1,296.
- Wade, J.A., 1977, Stratigraphy of Georges Bank Basin-- interpretation from seismic correlation to the western Scotian Shelf: Canadian Journal of Earth Science, v. 14, no. 10, p. 2274-2283.
- Wade, J.A., G.R. Campbell, R.M. Proctor, and G.C. Taylor, 1989, Petroleum Resources of the Scotian Shelf, Geological Survey of Canada Paper 88-19.
- Walper, J. L. and R. E. Miller, 1985, Tectonic evolution of Gulf Coast basins, *in* B. F. Perkins and G. B. Martin (eds.), Habitat of Oil and Gas, Program and Abstracts, Fourth Annual Research Conference, Gulf Coast Section: Austin, Society of Economic Paleontologists and Mineralogists Foundation, Earth Enterprises, p. 25-42.
- Waples, D. W., 1980, Time and temperature in petroleum formation--application of Lopatin's method to petroleum exploration: American Association of Petroleum Geologists Bulletin, v. 64, no. 6, p. 916-926.
- Weed, E. G. A., J. P. Minard, W. J. Perry, Jr., E. C. Rhodehamel, and E. I. Robbins, 1974, Generalized pre-Pleistocene geologic map of the northern United States Atlantic continental margin: U. S. Geological Survey Miscellaneous Geologic Investigations Map I-861, Scale 1:1,000,000.
- Williams, G. L., 1974, Dinoflagellate and spore stratigraphy of the Mesozoic-Cenozoic offshore Eastern Canada, *in* Offshore Geology of Eastern Canada: Geological Survey of Canada Paper 74-30, v. 2, p. 107-161.
- Williams, G. L., 1977, Dinocysts--their classification, Biostratigraphy, and paleoecology, *in* A. T. S. Ramsay (ed.), Oceanic Micropaleontology, v. 2, New York, Academic Press, p. 1,231-1,326.
- Williams, G. L. and W. W. Brideaux, 1975, Palynologic analyses of Upper Mesozoic and Cenozoic rocks of the Grand Banks, Atlantic Margin: Geological Survey of Canada Bulletin, v. 236, p. 1-163.

- Woollam, R. and J. B. Riding, 1983, Dinoflagellate cyst zonation of the English Jurassic: Institute of Geological Sciences Report, v. 83, No. 2, p. 1.
- Worsley, T. R., 1971, Calcareous nannofossil zonation of Upper Jurassic and Lower Cretaceous sediments from the Western Atlantic, *in* Proceedings II, International Planktonic Conference, Roma, 1970: Rome, Edizioni Tecnoscienza, p. 1,301-1,321.